SISMOGRAMMES HORS-CURRICULUM

Des Tsunamis aux Ondes de Gravité Aériennes

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TSUNAMI RECORDED ON SEISMOMETERS

- Horizontal long-period seismometers (GEOSCOPE, IRIS...) record ultra-long period oscillations following arrival of 2004 tsunami at nearby shores [*R. Kind*, 2005].
- Energy is mostly between 800 and 3000 seconds
- Amplitude of equivalent displacement is **centimetric**





[Hanson and Bowman, 2005]



CAN WE QUANTIFY SUCH RECORDS?



EXCITATION OF TSUNAMI in NORMAL MODE FORMALISM

• *Gilbert* [1970] has shown that the response of the Earth to a point source consisting of a single force **f** can be expressed as a summation over all of its normal modes

$$\mathbf{u}(r,t) = \sum_{N} \mathbf{s}_{n}(\mathbf{r}) \left(\mathbf{s}_{n}^{*}(\mathbf{r}_{s}) \cdot \mathbf{f}(\mathbf{r}_{s}) \right) \cdot \frac{1 - \cos \omega_{n} t \exp \left(-\omega_{n} t/2Q_{n} \right)}{\omega_{n}^{2}} \quad .$$

the *EXCITATION* of each mode being proportional to the *scalar* product of the force **f** by the eigen-displacement **s** at location **r**_s.

• Now, an *EARTHQUAKE* is represented by a system of forces called a *double – couple*:



The response of the Earth to an earthquake is thus

$$\mathbf{u}(r,t) = \sum_{N} \mathbf{s}_{n}(\mathbf{r}) \left(\boldsymbol{\varepsilon}_{n}^{*}(\mathbf{r}_{s}) : \boldsymbol{M}(\mathbf{r}_{s}) \right) \cdot \frac{1 - \cos \omega_{n} t \exp \left(-\omega_{n} t/2Q_{n}\right)}{\omega_{n}^{2}}$$

where the *EXCITATION* is the *scalar product* of the earthquake's **MOMENT** M with the local *eigenstrain* ε at the source \mathbf{r}_s .

This formula is directly applicable to the case of a tsunami represented by normal modes of the Earth.

QUANTIFYING the SEISMIC RECORD at CASY



Assume that seismic record (*e.g.*, at CASY) reflects response of seismometer to the *deformation of the ocean bottom*.

FORGET THE ISLAND (or continent) !

• Use *Gilbert*'s [1980] combination of displacement, tilt and gravity;

Apparent Horizontal Acceleration (Gilbert's [1980] Notation):

$$AV = \omega^2 V - r^{-1} L(gU + \Phi)$$

or (*Saito*'s [1967] notation):

$$y_3^{APP} = y_3 - \frac{1}{r \,\omega^2} \cdot (g \, y_1 - y_5)$$

• Use *Ward*'s [1980] normal mode formalism;

Evaluate Gilbert response on solid side of ocean floor, and derive equivalent spectral amplitude of surface displacement $y_1(\omega) = \eta(\omega)$.

- Use Okal and Titov's [2005] Tsunami Magnitude, inspired from Okal and Talandier's [1989] M_m ;
- Apply to CASY record at maximum spectral energy $(S(\omega) = 4000 \text{ cm}^*\text{s at } T = 800 \text{ s}).$

 \rightarrow Find $M_0 = 1.7 \times 10^{30} \, dyn - cm.$

Published: 1.15×10^{30} dyn*cm [Stein and Okal, 2005; Tsai et al., 2005]

Acceptable, given the extreme nature of the approximations.

 \rightarrow Suggests that the signal is just the expression of the horizontal deformation of the ocean floor, and that

CASY functions in a sense like an OBS !!

MAULE, Chile, 2010

8 Seismic Stations — 12 Components

 \rightarrow In the 500-2000 s period range, the results are generally in agreement with the CMT scalar moment.



This supports the finding [*Okal et al.*, 2010] *that the Maule earthquake is* **not a slow event**.

 \rightarrow At higher frequencies (not shown), the results would depend on the response of the individual island structure.

A REMARKABLE ANTECEDENT



Kuriles, 07 SEP 1918 at Apia [*ex*–German] Samoa

as reported by G. Angenheister [1920]

Beim Vorüberziehen der Flutwellen des Kurilenbebens (etwa 9^b nach Ankunft der seismischen Wellen in Apia) zeigte, der Wiechertsche Horizontal-Seismograph Neigungswellen von ¹/2^b-1^b Periode und bis zu 0".2 Amplitude.



2022 HUNGA-TONGA VOLCANIC EXPLOSION

A very intriguing "tsunami"

• At many locations of the Pacific, wave activity starts **BEFORE** the predicted arrival of the tsunami.

Time (mn after 05:00 GMT)

- → This corresponds to an acoustic wave in the atmosphere, which is coupled to the water column, resulting in a disturbance of the sea surface.
- That wave, propagating at a typical velocity of 313 m/s, is significantly precursory to the tsunami.

A SPECTACULAR EXAMPLE OF THE TWO WAVES

is given by the EW seismic record at Pitcairn Island

NOTE

- the perfect dispersion of the "true" tsunami, outside the Shallow-Water Approximation
- the weak dispersion of the air wave
- the strong spectral peak in the tail of the tsunami at ~4 mHz (stay tuned)
- the absence of conventional long-period seismic waves

Perhaps RATHER

AN OLD REVENANT...

AIR-SEA WAVES OBSERVED DURING 1883 KRAKATAU EXPLOSION

TSUNAMI GENERATION by Volcanic Explosions at Sea

Krakatau [Sunda Straits], 27 August 1883

ANAK KRAKATAU, Sept. 2016

Born 1927 ... and Still Growing !

A catastrophic tsunami killed 35,000 people in Batavia (Jakarta). *Nomambhoy and Satake* [1995] showed that it can be well modeled by an underwater explosion.

The tsunami was reported recorded world-wide (on tidal gauges), which would seem to contradict the dispersive nature of the short wavelengths associated with sources of small dimensions...

Press and Harkrider [1965, 1967] had shown that the tsunami is actually triggered by an **air wave** generated by an atmospheric explosion, and re-exciting the ocean as it propagates.

This explains

- the propagation of the "tsunami" along great circle paths occasionally crossing... a continent!
- the occasional early arrival of the tsunami at distant tidal stations (315 m/s as opposed to 200 m/s).
- and allows an estimate of the power of the explosion (100 to 150 Mt).

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When the atmosphere is underlain by an ocean, the two families of modes feature different boundary conditions

- with their eigenfunctions penetrating the other medium.
- \rightarrow But their dispersion remains [largely] unaffected.

EXCEPT in the range 0.15–0.25 mHz (400-550 s)

In this study, we will focus on the lower-frequency Frequency part of the spectrum, where the air-wave (GR_0) and tsunami modes are essentially uncoupled.

Additionally, at higher frequencies, selective resonance can take place where dispersion curves for S and GR modes intersect, also involving the spheroidal modes of the Solid Earth.

This can lead to largely monochromatic signals as observed at Pinatubo (1991) [Kanamori et al., 1994], and (remember) on the 2022 Pitcairn seismogram.

ANOTHER SPECTACULAR CASE OF AIR WAVES

"Tsar' Bomba", 30 OCT 1961

The largest ever nuclear test, 57 Mt

 14 ⁴ 11 ⁻ 18 ³ -11 ⁴	
 91 *35" (X(+ 17%)	
 	Longyearbyear Svalbard

Air wave recorded on a seismometer at Tsukuba, Japan (6100 km)

 \rightarrow No sea waves are known for this event...

Место взрыва

TSUNAMI

by

NEXT-DAY AIR ?

Now... Even Across Continents !

Of course, after crossing a continent the real tsunami disappears and only the air wave is left, including its multiple passages

HORIZONTAL PARTICLE MOTION at FDF (Geoscope)

First Passage of Air Wave

Particle motion is polarized $\sim 300^{\circ}$

about 45° from great circle (255°) !!

Polarization (if any...) varies greatly, occasionally over VERY SHORT distances (<100 km)

THEORETICAL COMPUTATIONS of AIR-[SEA]-WAVES

We use two codes

1. " HASH "

This code was written originally by D.G. Harkrider in the 1960s, following his theoretical work with F. Press.

It considers a flat-layered Solid Earth – [Ocean] – Atmosphere structure, and solves for the eigenfunction of the wave through an algorithm based on a Haskell propagator.

2. " MODE "

This code was written originally by H. Kanamori to compute the spheroidal (Rayleigh) modes of the Earth.

It traces its ancestry to the works of C. Pekeris, H. Takeuchi and their students, notably M. Saito [1967].

It was adapted to the case of tsunamis by EAO following the work of S. Ward in the 1980s, and to the case of air waves by P. Lognonné, E. Clévédé and H. Kanamori in the 1990s.

In this study, we simply added about 40 atmospheric layers from DGH's model, and found that the program computes the air [-sea] waves without the need for any significant changes.

AIR WAVE EIGENFUNCTION IS LARGELY INDEPENDENT OF OCEAN DEPTH

When defined as the overpressure y_2 , the structure of the air wave eigenfunction in the atmosphere is found to be largely independent of the presence (or depth) of the water column.

→ This simply expresses that the coupling with the water remains very weak and does not appreciably affect the structure of the wave (nor its celerity).

And then, because in a fluid layer, the normal mode excitation coefficient N_0 for an explosive source

depends ONLY ON THE COMPONENT y_2 ,

*N*₀ is ALSO INDEPENDENT OF THE DEPTH OF THE OCEAN (or, actually of its presence)

 \rightarrow The excitation of the air wave by an atmospheric explosion is not affected by the presence of an ocean below it.

WHAT is REALLY a *GR*⁰ AIR WAVE ?

Some back-of-the envelope thoughts based on simple Physics

 \rightarrow The "tsunami" of the atmosphere ?

In a sense, yes, BUT

- \rightarrow Such a tsunami should propagate (under the SWA) at a speed $C = \sqrt{g H}$
- But what is the height *H* of the atmosphere ? Probably, some average <*H* > weighted by the particle density which decreases fast with height.
- Various attempts to obtain such an average height yield anywhere from 9 to 15 km.
- Note that this number also gives a reasonable estimate of the atmospheric pressure at the surface of the Earth, for an average density of about 2/3 the surface density.
- This would fit very well a "tsunami" velocity of 313 m/s, requiring $\langle H \rangle = 10$ km under the SWA.

BUT....

A close examination of the eigenfunction of GR_0 shows that the energy is mostly... elastic. Strange ?

• We note that the speed of the "tsunami" (313 m/s) is VERY CLOSE to that of that of sound in the atmosphere (~ 340 m/s).

An interesting coincidence ?

→ Which brings in the question of the effect of finite compressibility on the structure and speed of a tsunami

This was investigated in the 2010s, notably by Watada et al., to explain early arrivals of tsunamis at teleseismic distances.

... but only for realistic values of the speed of sound in the ocean, which remains at least 7 times greater than that of a tsunami for all reasonable depths.

→ We investigate this effect in the case of a liquid 5-km deep ocean for compressional velocities ranging from 10 km/s (practically incompressible) to 0.1 km/s (less than the tsunami's SWA celerity).

We focus on the elastic fraction of the potential energy of the wave which should be close to 0 for a true "tsunami".

SPEED of SOUND α

CLOSE to CELERITY C !

Effect of *α* on oceanic tsunami

For realistic values of α, the elastic energy is at most 2%, no influence on structure of tsunami wave.

For α close to tsunami celerity, the energy becomes strongly elastic, the wave loses its structure as a tsunami.

313 m/s 340 m/s

An interesting coincidence ?

Maybe NOT.....

An interesting coincidence ? Maybe NOT....

• SWA Tsunami velocity

$$C = \sqrt{g < H} > \qquad \qquad C^2 = g < H >$$

• Sound velocity

 \rightarrow

$$\alpha = \sqrt{K_S / \rho}$$

For a **perfect gas**, $K_S = \gamma P$ ($\gamma = 1.4$ for N₂, O₂).

$$\frac{C^2}{\alpha^2} = \frac{g < H > \rho}{\gamma P}$$

In the relevant parts of the atmosphere, *P* is expected to be on the order of $\rho g < H >$, where < H > is an average value of the atmosphere thickness.

THUS, THE 2 VELOCITIES WILL BE COMPARABLE AND A PERFECT GAS CANNOT SUPPORT A TSUNAMI WHOSE ENERGY IS MOSTLY GRAVITATIONAL

MAJOR THEORETICAL QUESTIONS

- **1.** In the Air-Sea Wave, which are the parameters controlling the ratio of the air wave amplitude (essentially its pressure at the bottom of the atmosphere) to the amplitude of the sea-surface disturbance (as recorded by a maregraph)?
- → How will this "impedance" depend on the depth of the water layer (and of course, frequency) ?

NOTE that it has often been taken as

 $\rho_w g \approx 1 \, \text{mbar}/\text{cm}$

i.e., the hydrostatic value, for example correlating the underpressure at the eye of a hurricane with the amplitude of the static storm surge

WILL IT APPLY TO AN AIR-SEA WAVE ?

DYNAMIC RESPONSE RATIO

Results using D.G. Harkrider's 2–D code for flat-layered structures

on water depth

 \rightarrow They generalize *Harkrider and Press'* [1967] results, who had considered only a depth of 5 km, thus approaching a dynamic ratio of 1 cm/mbar, which also corresponds to the hydrostatic value

$$\frac{1}{\rho_w g} = \frac{1}{1.03 * 981} \frac{\text{cm}}{\text{barye}} = 0.99 \frac{\text{cm}}{\text{mbar}}$$

but this coincidence is an artifact of their choice of depth.

The decrease of the dynamic response with decreasing water depth predicts weaker coupling and smaller tsunami amplitudes in shallow basins, as exemplified in the Bering Sea during the Tonga explosion.

NOTE wave disappearing at Nome, on shore of very shallow Northern Bering Sea (< 200 m).

DIRECTIONS FOR FUTURE STUDY

• Understand lateral heterogeneity in air wave records

 \rightarrow Remember Professor Mohorovičić,

who switched careers from Meteorology to Seismology...

A. Mohorovičić 1857 – 1936

- Relate Seismic Moment (about 10²⁵ dyn*cm) to Energy (traditionally expressed in kt or Mt)
- Understand seismic recording of air wave by horizontal sensors
- **Revisit properties of Tsar' Bomba Air wave** (and search for any sea surface disturbance?)